1	SENSITIVITY ANALYSIS OF SEISMIC ATTRIBUTES PARAMETRIZATION TO REDUCE
2	MISINTERPRETATIONS: APPLICATIONS TO DEEPWATER CHANNEL COMPLEXES
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#### ABSTRACT

Geoscientists apply algorithms such as seismic attributes to better interpret depositional systems
that enhance various aspects of the seismic data. However, they are limited by the original seismic
amplitude or frequency content, data quality, and algorithm parameters considered.

Additionally, our capacity to interpret depositional system architecture is limited by seismic resolution, which results in potential misinterpretations associated with the correct position of stratigraphic features. This is particularly important as mapping reservoir architecture (geobody size, shape, and stacking patterns) in the subsurface is critical for exploring and producing hydrocarbons, CO<sub>2</sub> storage, and geothermal resource development since it can define connectivity or compartmentalization of flow zones.

To address these concerns, we investigated five synthetic seismic volumes from low to high-31 frequency bandwidths of 15 Hz, 30 Hz, 60 Hz, 90 Hz, and 180 Hz based on an architectural model 32 of an outcropping deepwater channelized slope system in the Magallanes Basin, Chile. We 33 analyzed 1) how seismic bandwidth affects the resolution of stacked stratigraphic features (i.e., 34 deepwater channel elements and Mass Transport Deposits (MTDs)) and their subsequent seismic 35 interpretation, and 2) the effect of different seismic attributes commonly employed in channel 36 interpretation on our data to understand the "mixing" or "vertical smearing" of stratigraphic 37 features by comparing the seismic with the true geological model 3) we explored how the 38 39 attributes' parametrization affects the imaging of differently sized features by modifying the 40 analysis window in each case from +/-2ms to +/- 50 ms. Finally, 4) we evaluated the effect of different noise levels in the sensitivity analysis. 41

Results show that the "mixing" of events occurs mainly as a result of 1) the seismic bandwidth, 2)
the algorithm used for each seismic attribute calculation, 3) the attribute vertical analysis window,

and 4) the signal-to-noise ratio of the data. Broadband, higher frequency data, and small analysis 44 windows provide clearer images of the stacked channels. In contrast, low-frequency data and larger 45 analysis windows result in more mixing or "composite" appearances, affecting interpretations and 46 net-to-gross estimates, especially in small-size stratigraphic features such as individual channel 47 elements and Mass Transport Deposits (MTDs). Our observations warn of potential 48 49 misinterpretations in applying default attributes to actual seismic data, especially in geometrical attributes and window-dependent ones. Recognizing these misinterpretations is paramount for 50 reconstructing deepwater architecture (this study), sedimentary and structural studies for drilling 51 locations, reserves estimation, and overall uncertainty assessment. 52

*Keywords*: deepwater; seismic facies; architectural elements; seismic geomorphology;
interpretation; seismic attributes; channel complex, analysis window.

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### 7 1. Motivation and objectives

Seismic exploration of deepwater channels is challenging due to the physical properties 58 inherent to the seismic, the variability in fill and stacking of reservoir geobodies, and the 59 uncertainty that can occur due to the lack of hard data (core or well data). Reservoir architecture 60 controls the distribution of fluids in the subsurface and the connectivity (or 61 compartmentalization) of the reservoir that impacts recovery or injectivity. Therefore, 62 geoscientists seek to understand how sensitive the interpretation of reservoir architecture is to 63 different quality and types of seismic data, as well as different attributes and parameters 64 commonly used to identify the architecture better and to make appropriate well plan decisions, 65 66 volumetric and recovery/storage estimates.

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The primary questions we aimed to address with the study are:

68	• How does the seismic data's frequency content affect the imaging of deepwater
69	architecture?

- What is the effect of each seismic attribute on the architecture interpretation and its true
  position, both vertically and horizontally?
- What is the effect of the seismic attribute analysis window size on the vertical smearing of
  architecture?
- What is the effect of seismic noise on our sensitivity analysis?

The study's importance is identifying common pitfalls in seismic interpretation using synthetic seismic data created from an outcrop-derived architectural model of a seismic-scale deepwater channel system. Studies with synthetic data like this allow geoscientists to understand uncertainty in interpreting channel architecture from seismic data.

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### 80 2. Introduction

When it comes to reservoir characterization using seismic reflection data, even if we 81 82 employ all the tools available to interpret, locate, and measure the reservoir that will contain (oil, gas, and water) or allow for the storage (geothermal, CO<sub>2</sub>) of economic resources, uncertainty 83 prevails. This interpretational ambiguity occurs due to changes in various physical parameters in 84 response to the media and the seismic records' inherent acquisition and processing characteristics. 85 The imaging and interpretation of different-sized stratigraphic features in the subsurface 86 87 using seismic reflection data are often compromised due to limits in seismic resolution, which in addition to tuning effects, can influence volumetric interpretations and gross rock volume 88 calculations (Pemberton et al., 2018). Therefore, geoscientists need to understand the common 89 90 pitfalls associated with seismic interpretation: the impact of the frequency content on the imaging

91 of reservoir architecture, the choice of parameters and attributes' influence on the interpretation of 92 architectural elements, and the detrimental impact noise can have on the overall picture. To address 93 this, we performed a sensitivity analysis that evaluated four parameters: frequency content, the 94 effect of the seismic attribute, the impact of the window of analysis, and noise level combined with 95 five 3D synthetic datasets.

The synthetic data that was derived from an outcrop analog in Magallanes Basin, Chile (Ruetten, 2021), and employed realistic acoustic impedances. The models that used a series of zero-phase Ormsby wavelets and 1D convolution (Langenkamp et al., 2021) allowed us to better understand how seismic bandwidth and seismic attribute parametrizations affect the resolution of stacked stratigraphic features in a seismic-scale channel system, providing insights that could be beneficial to the industry for drilling decisions, whether it is for hydrocarbon, geothermal, or CO<sub>2</sub> storage purposes.

In order to extract the most value and information from the seismic data, seismic 103 interpreters often derive seismic attributes from the data to reveal additional stratigraphic or 104 structural features. These attributes provide a means to enhance vertical and lateral changes in 105 reflectivity, thickness, continuity, and orientation of seismic features. From the exhaustive list of 106 107 seismic attributes existent, we focus on amplitude-derived, instantaneous, and geometric attributes for offering promising results in channel architecture definition (La Marca, 2020). All 108 109 coherence algorithms (that belong to the geometric attributes' class) use a vertical and lateral 110 analysis window, whether they are based on cross-correlation, semblance/variance, eigenstructure analysis, or the gradient structure tensor. For good quality data, Marfurt et al. 111 112 (1998) found it best to analyze stratigraphic features using a temporal analysis window as narrow 113 as possible, determined by the highest frequency in the data or the 3<sup>rd</sup> frequency corner in the

Ormsby wavelets. For poor-quality data, a larger window approximating the dominant period ofthe data provides improved results with minimal stratigraphic mixing.

Pemberton et al. (2018) and Langenkamp (2021) provided insights into the effect of amplitude and frequency on architectural element imaging and interpretation and facies classification. Nonetheless, the impact of seismic attributes, the parameters, and the noise content were not evaluated. Hence, this is one-of-a-kind study that focuses on assessing the complexities of attribute parameterization using synthetic data based on a known geologic model.

With the true model known, attribute parameterization and its effects on stratigraphic interpretation can be quantified, particularly highlighting the parameters that impact the apparent stratigraphic mixing or smearing of events such as the windows of analysis.

We first describe the aspects related to the architectural (outcrop) model and the characteristics of the synthetic datasets used. Then, our workflow is explained, providing details on the four parameters evaluated. Results are presented and focused on the cases derived from parameter combinations/sensitivity analysis. In the end, we provide a table and workflow that allows geoscientists to identify potential pitfalls in interpretation and address them according to their individual datasets, as best practices in interpretation should be documented and available to the geoscientific community to help reduce uncertainty in reservoir characterization.

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## 132 3. Geological model description

The geological model that is the basis of this study is derived from a sandstone-rich deepwater channel system along a progradational slope system (Hubbard et al., 2010). These deepwater slope deposits from the Late Cretaceous (70-80 Mya) Tres Pasos Formation are exposed on approximately 3 km long, 200m thick outcrops near Laguna Figueroa in the Magallanes Basin, Southern Chile (Macauley and Hubbard, 2013; and Hubbard et al., 2014). The
high quality of the stacked channel systems has been used to construct a seismic-scale 3D
architectural model of the deepwater channel system (Pemberton et al., 2018; Jackson et al.,
2019; Langenkamp et al., 2020; Ruetten, 2021). According to Fildani et al. (2013) the outcrop is
analogous to many slope channel systems globally in stratigraphy and depositional setting, which
makes it an excellent benchmark for any study that aims to address problems associated with
channel interpretation.

The models are the result of several studies from Macauley and Hubbard (2013), Fletcher 144 (2013), and Southern et al. (2017) combining measured sections, hierarchical stratigraphic 145 interpretations, paleoflow measurements and thousands of GPS data points that calibrated a drone-146 derived photomosaic. For these models, the fundamental architectural component are channel 147 elements, defined as distinct, mappable channelized sedimentary bodies (Figure 1). Multiple 148 stacked, related channel elements form a channel complex, and two or more complexes form a 149 channel complex set (McHargue et al., 2011; Macauley and Hubbard, 2013, Meirovitz et al., 2020; 150 Figure 1). The outcrops at Laguna Figueroa contain two complex sets, simply referred to as the 151 Upper and Lower Figueroa. The upper complex set consists of eight channel elements and are 152 153 grouped into four distinct channel complexes. The lower complex set contains twelve channel elements grouped into three channel complexes. Elements are modeled with a standardized width 154 155 of 400 m and thickness of 25 m. Three additional architectural components are present in the 156 outcrop: mudstone drapes at the base of channel elements, mass transport deposits (MTDs) at the base of channel complexes, and inner-levee thin-bed deposits encasing the channelized elements 157 158 (Macauley and Hubbard, 2013; Hubbard et al., 2014). The geological models consist of five facies: 159 1) channel element axis in yellow, 2) channel element off-axis in orange, 3) channel element 160 margin in brown, 4) homogeneous shale in gray, and 5) background shale (inner and outer levee161 facies) in white (Figure 2C).

Jackson et al. (2019) developed the first fine-scale geocellular model combining channel 162 planforms and vertical stacking for the lower outcrop section (lower channel system) but did not 163 include hierarchical groupings in the architecture. Pemberton et al., (2018) generated forward 164 165 seismic models using Jackson et al. (2018)'s model and analyzed seismic interpretation of architecture as a function of seismic resolution. Nielson (2018) analyzed the tuning effects of 166 single channel elements. Ruetten (2021) updated Jackson's initial model with new interpretations 167 and added an upper channel system separated from the lower system by a debris flow, and 168 studied how stacking patterns impact reservoir connectivity and fluid flow. Finally, Langenkamp 169 (2021) analyzed the influence of stacked channel element architecture on facies classification 170 using Ruetten's model. This work utilizes the geocellular model of Ruetten (2021) and synthetic 171 seismic models from Langenkamp (2021). 172

The five synthetic seismic models used in this study were built using a series of zero-173 phase Ormsby wavelets of 15 Hz, 30 Hz, 60 Hz, 90 Hz, and 180 Hz and 1D convolution (Chile 174 Slope Systems research consortium; Langenkamp et al., 202; Figure 2C) with a reflectivity 175 176 model. More aspects of each model are found in Langenkamp (2021). Facies-based rock properties (Figure 2B), adopted from Stright et al. (2014), show that amplitude peaks represent 177 178 an increase in acoustic impedance (Figure 2D). In contrast, troughs depict a decrease in acoustic 179 impedance. The synthetic volumes have a vertical window of 500 ms. For analysis purposes, we cropped the volume from 120ms to 380ms to avoid dead/blank zones in the reflectivity and focus 180 181 on the target channel systems.



Figure 1. Stratigraphic hierarchy: from a single element to complex sets. A vertical slice within the Pipeline 184 3D dataset offshore New Zealand shows the seismic appearance of different architectural hierarchies in 185 deepwater channels, with corresponding cartoons below. The smallest architecture (4th to 5th order) is 186 the channel element (box 1 in green). The second hierarchy (6<sup>th</sup> order) occurs when the channel elements 187 stack together, forming a channel complex (box 2 in blue). The higher-order hierarchy (7<sup>th</sup> order and higher) 188 189 occurs with the amalgamation of channel complexes, developing a channel complex set (box 3 in magenta). 190 The color legend indicates the distinct facies that commonly occur within each element of each architecture. 191 Measuring the sizes of each architectural element indicated on the right as well as their hierarchy, provides key insight into the underlying depositional processes as well as a prediction of the more common 192 lithologies. Hierarchies mentioned follow Pickering and Cantalejo (2015) classification. 193



195 Figure 2: A visual guide showing the steps to convert outcrop measurements to a synthetic model. (A) Location and 196 exposure of the outcropping deepwater channels at Laguna Figueroa. Paleoflow is from North to South (obliquely and 197 to the right into the outcrop at this location) (B) Conceptual diagram of the Upper and Lower Figueroa outcrops 198 showing channel elements, complexes, and complex sets. The red line indicates the outcrop profile. Left of the line is 199 into the outcrop face, and the right has been eroded away. (C) geocellular model using the constraints from (B) 200 augmented by facies and corresponding rock properties, including acoustic impedances from Shallow Offshore West 201 African modeled rock properties (Stright et al., 2014). (D) The Ormsby wavelet and a representative vertical slice 202 through the 3D synthetic seismic data volume generated from the model shown in (C). Courtesy of Teresa 203 Langenkamp and Lisa Stright.

204 **4.** Methods

To address the questions posed in the study and focus on analyzing the effect of 205 bandwidth on vertical resolution, we introduce the term "stratigraphic event mixing." This 206 vertical smearing phenomenon is explained in Figure 3, using a real example where the 207 interference of channels from other stratigraphic levels is evident. As depicted in Figure 4, first, 208 209 we performed an exploratory data analysis to define a vertical window of interest constrained to the objective of the study: 120 ms and 380 ms from our five volumes of synthetic seismic data. 210 After cropping the 15 Hz, 30 Hz, 60 Hz, 90 Hz, and 180 Hz dominant frequency volumes, we 211 performed a sensitivity analysis on four parameters: 1) frequency content, 2) seismic attribute 212 effect, 3) the effect of the analysis window, and 4) the influence of band-limited random noise. 213

Next, we calculated a series of seismic attributes from amplitude accentuating, geometric, 214 and instantaneous attributes commonly employed in seismic interpretation of channel systems. 215 Due to the number of cases to evaluate, we decided to explore and present the most 216 217 representative seismic attributes for each case in detail (the most commonly used and that provided better results). Next, for each scenario, we defined a suite of 3-trace by 3-trace analysis 218 windows with various vertical lengths from 2 ms to 50 ms. For visualization and interpretation 219 220 purposes, we used co-rendering techniques. Finally, we explored the impact of the addition of low and high levels of noise. 221

The final analysis of the results was performed by combining the aforementioned parameters into the following cases: 1) the response on the same attribute and analysis window in the different bandwidth volumes, 2) the impact of changing the analysis window for different seismic attributes, 3) the effect of changing the window of analysis size for the same seismic attribute, and 4) the effect of noise. All cases were contrasted with the initial actual data/model.







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# 243 4.1. Parameter 1: frequency content effect

Tuning thickness is the temporal resolving power of seismic data. Some authors use resolution and tuning thickness terms interchangeably, although tuning starts right below the vertical resolution. Table 1 presents the tuning thickness for each element in the synthetic volumes studied here. Knowing the resolvability in seismic is paramount to understanding the effect of other parameters considered in the sensitivity analysis.

Resolution is the ability to resolve by seismic interpretation methods two features that are close together. By definition, the vertical resolution of seismic data is <sup>1</sup>/<sub>4</sub> of the wavelength ( $\lambda$ ), where the  $\lambda$  is determined by dividing the average velocity by the dominant frequency.



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Figure 4. Workflow of the study. Four different parameters were evaluated in the sensitivity analysis: 1) the effect of frequency content, 2) the impact on the choice of the seismic attribute, 3) the analysis window effect, and 4) the sensitivity to band-limited random noise addition. The best cases were selected to be shown in each case, and analysis was performed by comparing them with the original 3D geological model derived from the Laguna Figueroa Deepwater outcrop. This resulted in 4 cases for analysis, ultimately leading to a workflow and documentation of best practices in channel architecture interpretation. The frequency content was the first parameter evaluated for the impact on the interpretation and resolvability of the different architectural elements. The five dominant frequency synthetic seismic volumes were analyzed, and we compared the results with those of

Pemberton et al. (2018), and Langenkamp (2021).

	Channel elements	MTDs		
Dominant Frequency (Hz)	Tuning Thickness (m)	Tuning Thickness (m)		
15	48.9	53.2		
30	24.4	26.6		
60	12.2	13.3		
90	8.1	8.9		
180	4.1	4.4		

Table 1. Tuning thicknesses for shallow and deep elements in each synthetic seismic volume. Modified
from Langenkamp (2021).

# 265 4.2. Parameter 2: Choice of seismic attributes

266 A seismic attribute is a computation made from algorithms applied to seismic data to get a more interpretable output. These responses relate to rock physical properties (La Marca, 2020) 267 in rocks and fluids in the subsurface. However, there are tens if not hundreds of seismic 268 269 attributes (Barnes, 2016), and time constraints do not allow for testing them all. Some of the latest studies (Posamentier and Kolla, 2003; Chopra and Marfurt, 2007; Hossain, 2020; and La 270 Marca and Bedle, 2022) have proven the successful application of amplitude accentuating, 271 geometrical and instantaneous attributes applied to PSTM data to interpret and characterize 272 channel elements and complexes in both fluvial and deepwater settings. Generally, we need the 273 combination of a geometrical attribute that allows defining edges and at least one attribute that 274

- provides insights into stratigraphy (La Marca et al., 2019) to characterize channel features in
- seismic data. Therefore, we focused on testing attributes that belonged to those three attribute
- 277 categories and chose the most prominent of each class.
- 278 Table 2 summarizes the most representative seismic attributes selected for each class: root mean
- square amplitude (RMS), Sobel filter coherency, and instantaneous frequency.

Attribute	Appearance	Attribute Category	Measurement	Use in architecture interpretation	
Coherence	a	Geometric	Direct measure of waveform similarity or how similar waveforms or traces are in a volume- used to emphasize continuous events or edges.	Delineates edges of channel elements	
RMS amplitude		Amplitude/Energy	Measures the square root of the average energy within a vertical window.	Provides measure of channel element vs inner levee Provides statistical measures of channel element fill between two picked horizons	
Instantaneous frequency	High Low	Spectral	A simple approximation of the mean frequency of the seismic wavelet.	Channel element thickness	

- Table 2. Selected seismic attributes to perform the sensitivity analysis showing a representative image,
  feature measurement, and use in channel element interpretation.
- 283 *4.2.1. RMS amplitude*
- 284 The RMS amplitude is an amplitude accentuating attribute often used for stratigraphic
- and lithologic variations enhancement. It is defined by the standard deviation,  $\sigma(t)$ , of the data,
- 286 d(t), within a running analysis window, subsequently measuring the reflectivity within that
- window (Meek, 2015). For a window that ranges from  $-T = -K\Delta t$  to  $+T = +K\Delta t$  about a sample j,
- the RMS amplitude is:

289 
$$d_{RMS}(j\Delta t) = \sigma(j\Delta t) = \left(\frac{1}{2K+1}\sum_{K=K}^{+K} \{d[(j+k)\Delta t)]\}^2\right)^{1/2} \quad . \text{ Eq } (1)$$

In this study, we evaluated how well RMS showed stratigraphic variations.

# 291 *4.2.2. Coherence (Sobel Filter)*

292	Seismic coherency is a measure of how similar traces are among their neighbors, which is					
293	a response to lateral changes in the seismic record caused by variations in structure, stratigraphy,					
294	lithology, porosity, and the presence of hydrocarbons (Marfurt et al., 1998), and it is determined					
295	computing amplitude derivatives along structural dip. Sobel filter (Luo et al., 1996) is one of the					
296	many coherence methods, which for seismic data normalizes coherence data to produce results					
297	between 0 and 1, where 0 is the lowest coherence, and 1 is the highest coherence. It has proven					
298	to be effective in delineating channel element edges (La Marca and Bedle, 2021; Hossain, 2020;					
299	and Herron, 2011); therefore, we aimed to test the definition of the channel elements' edges.					
300	4.2.3. Instantaneous frequency					
301	The instantaneous frequency is computed sample by sample, is the time derivative of the					
302	instantaneous phase $\varphi(t)$ :					
303	$F(t)=d[\varphi(t)]/dt.$ Eq (2)					
304	and provides a simple estimate of the mean frequency of an isolated seismic event.					
305	Subrahmanyam and Rao (2008) find that the instantaneous frequency attribute can					
306	indicate bed thickness and provide lithology insights. Chopra and Marfurt (2007) emphasize					
307	their usefulness in identifying abnormal attenuation and thin-bed tuning. Our model has thin					
308	features, such as channel elements (from axis, off-axis, to margins), that we aimed to test by					
309	using this instantaneous attribute.					

### 310 4.3. Parameter 3: Analysis window effect

According to Lin et al. (2014), the scale of the window height, *H*, is a function of the

312 dominant or peak frequency  $f_{\text{peak}}$ :

313 
$$H=1/(2f_{\text{peak}})$$
 . Eq (3)

314 Applying this concept to our dataset spectrum, we would need to use a window of 2ms

for the highest frequency volume (180Hz) and ~33 ms for the lower frequency volume (15Hz).



Figure 5. A representative analysis window used in attribute calculations. Signals are sampled at discrete points, not continuous recordings. Therefore, each seismic trace will contain as many samples as the sample rate allows. For the example shown, if we consider the dominant period (distance between two peaks), our analysis window will contain 20 samples, equivalent to ~ 20 msec. Other examples of analysis windows are depicted, including the smallest possible equivalent to the sample increment, in this case, 1 msec.

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To understand the concept of analysis window, we have drawn a cartoon (Figure 5) that

- shows a seismic trace with its respective samples. It illustrates how a small, medium or large
- analysis window would look and what a dominant period is. Usually, the number of samples
- 325 considered in a default analysis window parameter setting is around 11 samples (default window
- 326 60ms) for a 6ms sample rate dataset; in our dataset, the sample rate is 1ms. In real data,
- 327 shallower, higher frequency data often shows smaller periods than the deeper strata/intervals

where lobes of the traces are wider. Due to this change in frequency, some authors, like Lin et al. 328 (2014), recommend using an adaptive window. However, this is not often possible or available in 329 the software packages commonly used by geoscientists. Nonetheless, the geoscientist/interpreter 330 can control the analysis window in the attribute settings. Therefore, we want to provide insights 331 into the difference between a default analysis window and a shorter or larger one (which reduces 332 333 or increases the number of samples, respectively). We compared the results from each different analysis window to the original, true data (the geological model), which provides a good 334 sensitivity analysis for all kinds of datasets and any of the attribute families studied here. 335

# 336 4.4. Parameter 4: Addition of band-limited random noise

When seismic data is recorded, we find two components: signal and noise. The latter 337 comprises all the unwanted recorded energy that contaminates seismic data, and it can be random 338 or coherent (Kumar and Ahmed, 2021). Random noise is generated by activities in the 339 environment where seismic acquisition work is being carried out, and this noise appears in a 340 seismic record as spikes (Enwenode, 2014). Seismic noise levels depend on the type of 341 acquisition—land or marine—and the intrinsic conditions unique to every site, such as climate, 342 the burial of sensor, and wind (Tanimoto et al., 2015). Although there are many types of noise, 343 like Gaussian, Pink, Brownian, violet, and blue, in this study, our fourth parameter incorporated 344 in the sensitivity analysis is the band-limited random noise. 345

Signal-to-noise ratio (SNR) is a measure used to compare the degree of signal to the level of background noise, in which case a ratio larger than 1:1 suggests more signal than noise. So, the lower the SNR, the noisier our dataset. As Chen et al. (2019) mentioned, this will lower the quality of the seismic, affecting subsequent analyses such as imaging and inversion.

350	We incorporated different noise levels, from a low noise added of 5% to a high 30%.				
351	Higher noise levels were not presented due to the incapacity to extract meaningful interpretations				
352	from the data. Nowadays, noise can be added to synthetic datasets thanks to available software				
353	like the one used in this study.				
354	5. Results				
355	To address the study questions and link the parameters taken into account in the				
356	sensitivity analysis, we present the results summarized in four cases:				
357	5.1. Case 1: The effect of the spectral bandwidth on the imaging of architectural elements				
358	After evaluating the five synthetic models of 15 Hz, 30 Hz, 60 Hz, 90 Hz, and 180 Hz				
359	dominant frequencies, we observed that the level of detail of the different sized architectural				
360	elements increases with frequency; therefore, broadband, higher frequency provides better				
361	resolution.				
362	Figure 6 shows the effect of each frequency in imaging each geological element. Figure				
363	6B corresponds to the lowest frequency, and we can distinguish channel complexes and mass				
364	transport deposits (MTDs). However, smaller features like the 25m thick channel elements and				
365	MTDs are mixed in thicker, unresolvable reflectors.				
366	Figures 6C and 6D show frequencies commonly encountered in the subsurface (30Hz and				
367	60Hz for vintage and recent data, respectively). Here, the individual complexes are well-defined.				
368	Nonetheless, the amalgamation or stacking of elements presents a single response and minimal				
369	acoustic impedance contrast.				
370	Higher frequencies shown in Figures 6E and 6F present the best responses compared to				

371 the original model. Figure 6E can even differentiate some of the stacked packages; overall, the

individual and stacked channel elements can be better resolved. It is noteworthy how the highest
frequency (180 Hz) illustrated in Figure 6F starts to lose the definition of the inner reflectors.
Another observation is that, in all cases, the vertical channel axis (center) is resolved better than
the channel's margins, both as a function of thickness (channel elements are thickest in the center
at the axis) and acoustic impedance contrast (rock properties of channel element margin is more
similar to inner levee than the channel element axis).

In general, channel complexes sets are visible at all frequencies analyzed, whereas individual complexes start to be resolvable from 30 Hz and higher. However, when elements have vertical stacking, channels do not show contrast in acoustic impedance due to repeated material/ similar composition and properties, therefore the attribute response is also affected by this phenomenon.

## 383 5.2. Case 2: Attribute sensitivity to thickness and extent of stratigraphic elements

After performing the seismic attribute sensitivity analysis, we noticed that the thickness and extent of stratigraphic events imaging are inherently linked to the frequency content of the seismic, analysis window, and seismic attribute used. At peak frequencies commonly encountered in the subsurface (around 30- 60Hz), the number of complexes was underestimated, and the size, shape, and type of architectural bodies (channels elements vs. margins, elements vs. complexes) were difficult to differentiate (Figure 7).

- 390 In most cases, channel complex sets were able to be interpreted. However, the
- amalgamation of smaller-scale channel elements results in an incorrect estimation of thicknesses.

392 Nielson (2016) documented the same phenomenon.



**Figure 6.** A representative vertical slice (224) and time slices at t=-191ms) through the (A) 3D model that shows (B) 15 Hz data, (C) 60 Hz data, (D) 90 Hz data, and (E) 180 Hz data. Notice the improvement in the channel architecture's detail with the increase in the dominant frequency and corresponding spectral bandwidth, being able to interpret complexes from figure B on.

For example, in cases where similar facies were in contact, the acoustic impedance
similarities did not allow for the differentiation of individual channel elements. Also, in tests
performed with a large analysis window, the interpretation of smaller features was not possible
(Figure 8B, 8D).



Figure 7. Results on the attribute sensitivity analysis to thickness and extent of architectural elements using
a default analysis window. We present results on the 30 Hz dataset, analogous to vintage seismic data (to
the left), and the 60Hz dataset, representing modern seismic data (to the right). Each attribute is presented
for both frequencies in a representative inline (224) and time slices (-280 ms and -191ms) as follows: (A)
RMS (Root Mean Square) amplitude for 30 Hz (B) coherence (Sobel filter) for 30Hz, (C) Instantaneous
frequency for 30 Hz, (D) RMS amplitude for 60 Hz (E) coherence (Sobel filter) for 60Hz, (F) Instantaneous
frequency for 60 Hz.

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# 411 5.3. Case 3: Effect of the analysis window size on the vertical smearing of stratigraphic

412 architecture

413 To demonstrate the effect of the window size on the vertical smearing of the different

- architectural elements in the model, we show an RMS amplitude sensitivity analysis (Figure 8).
- 415 We perceived that RMS amplitude offers a detailed picture of the various facies in the 180 Hz
- 416 volume, including imaging of the MTD associated with the channel complexes and the

individual channel elements. However, for larger analysis windows and lower frequency 417 volumes, for example, 50 ms window combined with a 15Hz dominant period seismic volume, 418 such details are lost, resulting in an accentuated vertical stratigraphic smearing effect. Therefore, 419 the application of amplitude-derived seismic attributes results more effective in higher frequency 420 content datasets and using a small analysis window than in another configuration. It is 421 422 noteworthy that a higher value of RMS is presented where an amalgamation of events occurs, which potentially leads to the interpretation of higher NTG (Figures 7 and 8). 423 Since the Sobel filter coherence seismic attribute aids in detecting discontinuities, such as 424 geological structures and edges, especially in time slices, we included this attribute in our 425 analysis to identify the channel complexes and the edges of the channel elements (Figure 9). 426 We observe stratigraphic mixing occurring in the vertical slices when we utilize a larger 427 analysis window. It is also noticeable that higher dominant frequency data in combination with a 428 small analysis window allows for a correct placing of feature edges (Figure 9C), which is 429 supported when compared to the true model. In contrast, the position of the channel element 430 edges deviates from the original/ true location or becomes distorted when the analysis window 431 size increases (Figure 9D) or the frequency of the data is small (Figure 9A) regardless of the 432

433 analysis windows used.

# 434 5.4. Case 4: Effect of noise in the sensitivity analysis

Figure 10 shows the most prominent results of the sensitivity analysis on the noise effect. We
compared, in this case, how instantaneous frequency, coherence, and RMS amplitude seismic
attributes respond to variations in band-limited random noise from 5% to 30%.



Figure 8. RMS amplitude sensitivity analysis to evaluate the effect of the window size W on the vertical smearing of the different channel elements in the model. (A) 15 Hz with W=2ms (B)50 Hz with W=50ms(C) 30Hz with W=2ms (D) 50 Hz with W=50ms (E) 180 Hz with W=2ms (F) 180 Hz with W=50 ms. Highfrequency data with a small analysis window provides the most suitable representation of the true model. In RMS amplitude, the most accurate facies depiction is given by high frequencies and small window combination, imaging channel element base, and inner and outer levee facies.



Figure 9. Sobel filter coherence attribute sensitivity analysis to evaluate the effect of the window size Won the vertical smearing and displacement of the different architectural elements in the model. (A) 15 Hz with W=2ms (B)30 Hz with W=2ms (C) 60Hz with W=2ms (D) 60 Hz with W=50ms. Notice how at a higher dominant frequency (60Hz) and smaller window of analysis, there is better detection of the channel element edges and channel element fill. In contrast, at lower frequencies and or larger windows of analysis, there is interference from deeper channel elements, as in the case presented on Figure 3.





455 particular result found is a low coherence value predominance when evaluating the 30% noise

456 case (Figure 10B).

## 457 5.5. Pitfalls in attribute interpretation and how to minimize them

- The results mentioned above warn of potential pitfalls when interpreting seismic data. For example, not understanding the frequency content and quality of the seismic dataset could result in the inappropriate use of seismic attribute parameters and subsequent misinterpretations. To summarize all these results, we present a list of common pitfalls (Table 3) in attribute interpretation
- and how the interpreter can attempt to avoid them.

Common	How to address it	Comments	
misinterpretation			
Not inspecting the frequency content and dominant frequency of the data	Inspect the seismic volume and understand dominant frequency and variations with depth	Higher frequency, broader bandwidth data provide higher vertical and lateral resolution	
Not understanding the quality of the data and resolution	Calculate the seismic resolution	Imaging of different scales of architecture improves with a higher signal-noise ratio	
Using meaningless attributes	Have a clear geological goal and calculate attributes effective for similar targets in literature	Avoid attributes that make pretty pictures but are not directly related to the target objective.	
Using default parameters	Select parameters adequate to your study. In windows-based attributes, smaller windows provide less vertical mixing whereas larger windows are less sensitive to noise. Because data quality and resolution change with depth, parameters appropriate for the shallow part may be suboptimum for the deeper part of your survey.	Defaults are provided for the most common cases encountered by the most common user (e.g., oil and gas exploration). If your data are unique in either acquisition or objective, test a wide variety of parameters and choose the ones that best delineate your target.	
Interpreting data that has a low S/N ratio, or high noise	Inspect your seismic volume and categorize the noise type and level when possible. Then apply a noise reduction or removal algorithm	Noise can hinder the interpretation of seismic facies and calculating attributes won't help. If data are noisy, we recommend just using amplitude- related attributes.	

463 **Table 3**. Common pitfalls in seismic interpretation and how to avoid them.



Figure 10. Sensitivity analysis of the effect of random noise in channel architecture interpretation. Results shown refer to the 60 Hz dominant frequency seismic volume (which would be analog to real datasets). (A) 5% band-limited random noise added evaluated in RMS, Coherence, and Instantaneous frequency attributes. Interpretation of prominent features like channel complexes is only possible. (B) 5% bandlimited random noise added applied to RMS, coherence, and instantaneous frequency attributes. Notice how the increase in noise is detrimental in the channel architecture interpretation, especially when using coherence. This may be due to the sensitivity that small windows have on high noise content.

#### 472 **6. Discussion**

### 473 6.1. The importance of mapping channel architectures

474 Channels systems, whether fluvial or deepwater in origin, exhibit petrophysical
475 characteristics that make them excellent reservoirs for oil, gas, or other resources (Slatt et al.,
476 2009). However, channel system mapping in the subsurface becomes challenging, especially
477 when only seismic data is available.

In seismic reflection data, the amplitude, frequency, and noise content will condition the quality of the seismic image, affecting the interpretation of different architecture and facies embedded in the reflector configuration and overall seismic response. It is, therefore, paramount to understand how size-dependent architecture is displayed and imaged under different conditions and how the application of seismic attributes could help or hinder the interpretation of such architecture.

In this study, we used a unique approach by employing 3D synthetic seismic datasets as a benchmark to perform a sensitivity analysis to understand how different-scales of architecture appear as a function of frequency content, noise level, type of seismic attribute, and parametrization, especially the analysis window selected to calculate attributes.

The analysis of all the scenarios resulted in the following observations: broadband higher frequency data (e.g., 90 and 180 Hz) combined with a short analysis window (e.g., 2 ms, 20 ms) minimized the stratigraphic mixing (Figure 6). In contrast, lower frequency data that were analyzed using a large vertical analysis window (Figure 8) resulted in poor imaging of the channel architecture, vertically mixing stratigraphic architecture at different hierarchical levels. This affects the temporal evaluation of features that show an overlap of individual architecture 494 (i.e., channel elements) in the system in consecutive time slices, or in other words, a vertical
495 offset from the known position (Figure 9) of sedimentary units (also shown by Pemberton et al.,
496 2018).

497 This observation warns of potential interpretation pitfalls in applying such seismic 498 attributes to actual seismic data. Stratigraphic mixing can hinder the correct temporal and spatial 499 representation of individual channel elements and boundaries of channel complexes, leaving the 500 internal architectures and potential fluid flow barriers imprecisely imaged (Coleman, 2000), 501 hence, incorrectly mapping and estimating the volume of the reservoir units of interest, which 502 could result in important economic loses.

# 503 6.2. The effect of frequency content in the imaging of architectural elements

Resolution has always been one of the main concerns for seismic interpreters since important features like small channels, or DHIs/ bright spots can be overlooked in seismic data. Also, hazards and baffles (Cardona, 2020; Meirovitz et al., 2020; Ruetten, 2021), like the MTDs, may not be imaged in the seismic picture analyzed. In this study, MTDs are only five meters thick and usually mantle the base of the channel complexes and complex sets.

To improve the seismic data's resolution, as demonstrated in this study by the enhanced definition of each element in higher frequency datasets, frequency content should be increased. This suggests that ideally, modifying the initial design of the seismic data and obtaining higher frequency data will allow us to get the resolution required to image submarine channel complexity to the detail commonly observed in outcrop (Coleman et al., 2000). This type of data is, however, costly. Although the cost of high-quality data is currently high, with the advancement in technology, we presume that access to higher-quality data at reduced prices will become available soon. An example of this kind of data is the use of OBN (ocean bottom nodes) – highresolution 3D seismic acquired with the addition of P-Cable, an offshore seismic data acquisition system that provides highly detailed ultrahigh-resolution images of the seafloor and subsurface geology (McGregor et al., 2022) – and high-resolution (1-6 kHz) chirp data that offer a better image of the stratigraphy and structure of rocks in the shallow subsurface.

It is likely that soon most of the seismic data acquired will be high resolution, without necessarily needing to be high frequency, and that we will find ways to improve our algorithms to treat the noise associated with acquisition and processing or overall improve the data quality from the early stages of seismic acquisition. But, until then, we need to understand the limitations of various kinds of seismic data and become aware of potential pitfalls in interpretation.

528 6.3. Selecting seismic attributes and parameters that are ideal for mapping channel
529 architecture

In this study, we evaluated the effect of seismic attributes on the channel system
architecture interpretation and their true position in the outcrop model. Our first insight is that
there is no one-size-fits-all kind of seismic data.

Therefore, we suggest that the first step when interpreting seismic reflection data should be to 1) define the geological goal, 2) become familiar with the acquisition and processing of information, 3) make an initial inspection of the data, 4) determine if there is some noise or artifact that the interpreter needs to correct or be aware of. Finally, 5) select seismic attributesbased on established purpose/geological goal.

538 One of the most critical steps is defining a clear geological goal. The interpreter should 539 reflect on what attribute would better suit their purpose. For example, we would use amplitude-540 derived or instantaneous attributes in studying bright spots. For channel systems, instantaneous 541 or frequency attributes can help highlight the differences between the channel elements 542 (Fedorova, 2016), whereas geometrical attributes (such as Sobel filter coherence) aid in 543 delineating the external shape of the architecture (La Marca, 2020).

Because there are tens to hundreds of attributes, testing many of them can be timeconsuming. Imagine that in this study, with just three attributes shown and the case combinations, we had a total of 90 volumes to evaluate. Therefore, it is recommended to work only with attributes whose principles the interpreter understands or, if using multiple attributes, rely on experimental designs like Box Behnken (Ferreira et al., 2007) that will help to synthesize a large amount of data.

Attribute computation varies from software to software, and some attributes are computed trace by trace, whereas others are sample by sample. Understanding this initially would help in setting the correct parameters in each case. For attributes that are window based, like wavelet or average frequency, and the ones studied here, the variation of the analysis window affects the imaging of architectural elements.

We used instantaneous frequency since it can indicate the edges of thin low-impedance thin features. Additionally, it is an excellent bed thickness indicator, where higher frequencies indicate sharp interfaces or thinly bedded strata and lower frequencies indicate thickly bedded sandstone-rich strata (Subrahmanyam and Rao, 2008). Interestingly, in the study, the true

thickness of isolated channel elements (approximately 25m of thick-bedded amalgamated
sandstone at the channel element axis) was more interpretable from the dataset of 60 Hz
dominant frequency and above and only partially distinguishable in the shallow portion of the
30Hz dataset. This finding was also noted by Nielson (2018) and Langenkamp (2021), although
neither study included seismic attribute analysis.

RMS amplitude measures reflectivity within a time window (Meek, 2015). It computes 564 the square root of the sum of squared amplitudes divided by the number of samples within the 565 566 windows used (Equation 1). Therefore, the number of samples and windows used affects the strength of reflectivity we get with the algorithm used. In contrast, some attributes like Sobel 567 filter coherence do not necessarily improve channel architecture imaging by selecting a minimal 568 analysis window. In this case, since the wavelength increases with increasing velocity, which 569 increases with depth, we agree with Lin et al. (2014) that coherence attributes should use a 570 shorter analysis window in the shallow section and a larger vertical analysis window in the 571 572 deeper section. Some software has this already integrated as an adaptive window, which we consider to be one route to improve attribute results in the future. 573

## 574 6.4. Impact of noise on the interpretation of channel features

Different types of noise could be found in our seismic data: coherent noise, a series of unwanted signals that appear when the source is applied (Alderton and Elias, 2021), and incoherent noise, which would appear whether we shoot or not. For this sensitivity analysis, we only explored the effect of band-limited random noise, which belongs to the latter category, on imaging the different architecture in a deepwater channel system.

580 When using instantaneous attributes such as instantaneous frequency, it was noticed that 581 at low noise levels (e.g., 5%), smaller features like MTDs were still visible. However, with the

increase in noise, it was extremely difficult to interpret individual geologic features in theseismic volume.

584 As Herron (2011) stated, the output's quality of interpretability relies on the input's noise content. In fact, a highly noisy dataset will likely contain very little reliable information. 585 Moreover, we noticed that the effect of stratigraphic mixing is emphasized in the coherence 586 587 attribute, where anomalous low or high values of the attribute were unexpectedly found and that the interpretation of small or large size features in the seismic was very hard to impossible to 588 perform. These findings stress the importance of noise removal using adequate techniques 589 related to the type of noise. This must be done in the early stages of seismic interpretation to 590 avoid misinterpretations (Figure 11). 591

In terms of the analysis window selected in the presence of noise, using a very large dataanalysis window (in three dimensions) will include plentiful data and likely produce output with a marked structural overprint. On the contrary, a window that is too small will barely include data and produce an outcome that is more a manifestation of noise in the data rather than geological content (Herron, 2011).

597 Our final thoughts are that although studies that use synthetic data and perform sensitivity 598 analyses provide a tool to address misinterpretations encountered in the seismic interpretation of 599 potential reservoirs, especially in the exploration stage, it is necessary to take into account that 600 each seismic dataset and geologic setting are unique to the exploration area and that every aspect 601 should be considered carefully before making impactful decisions.



Figure 11. Generalized workflow for a geoscientist to avoid pitfalls in interpretation by getting optimizedresults according to their dataset.

609 7. Conclusions

The sensitivity analysis of 3D synthetic seismic volumes derived from a model of an outcropping deepwater slope-system and assigned reservoir properties was performed. The study comprised the combination of four parameters: 1) frequency content, 2) attribute selection, 3) windows of analysis and 4) noise content to identify their effects on the imaging of different scales of deepwater architecture.

The results are summarized in four cases that allow the depiction of common pitfalls in channel interpretation: 1) the effect of the spectral bandwidth on the imaging of the different scales of architectural elements, 2) the seismic attribute sensitivity to the thickness and extent of element and complexes, 3) the effect of the analysis window size on the vertical smearing of channel elements and MTDs, and 4) the effect of noise in the sensitivity analysis.

In this study, we introduced the term "stratigraphic mixing" to define the combined picture resulting from the inability to resolve a single channel element. In this sense, broadband, higher bandwidth (e.g., 90 Hz) data combined with a short analysis window (e.g., 2ms) minimizes stratigraphic mixing. In contrast, lower bandwidth data, in addition to a large analysis window, results in poor imaging of the channel element and channel complexes that exhibit a "composite" appearance, vertically mixing geological features at different stratigraphic levels.

The importance of this analysis resides in that stratigraphic mixing affects the temporal evaluation of features that show an overlap of individual architecture at different scales in the system in consecutive time slices or a vertical offset from the known position of sedimentary units, which may result in important economic losses by misplacing an exploration well (e.g., actual target not in place) or overestimating reserves due to an incorrect NTG estimation and subsequent volumetric calculation (e.g., baffles like MTD are not imaged). In order to determine the sensitivity of the different architectural elements to seismic
attributes, we explored three classes of algorithms: amplitude-based, geometrical, and
instantaneous, and showed the results of the most prominent attributes: RMS amplitude,
instantaneous frequency, and Sobel filter coherence.

Attributes that are window-dependent, such as RMS amplitude, show an improved imaging of the actual thickness of the channel architecture when calculated using a short analysis window (e.g., 2ms) over a higher frequency dataset (from 30Hz and above). With larger analysis windows and or a small frequency dataset, there is significant vertical stratigraphic mixing. High RMS values and a composite effect were found in stacked channel element configurations indistinctly of the parameters used, which would likely result in an overestimation of NTG.

642 Conversely, for edge detection attributes like coherence, a small analysis window does 643 not provide a better depiction of channel element/complex/complex set edges. Instead, we 644 observed displacement with respect to the actual position and composite pictures of them, 645 especially in vertical sections, which makes their interpretation cumbersome. Therefore, we 646 suggest using an adaptative window with depth or a default analysis window (half of the peak 647 frequency is a good approximation).

648 When evaluating the last parameter, which corresponds to the effect of band-limited 649 noise content, it was observed that the mapping of channel elements is hindered by adding noise 650 to the data. When the noise level increases, as expected, the interpretation of features is hindered 651 by the impact of the noise. We suggest applying algorithms that will allow us to eliminate or 652 reduce the noise before calculating any seismic attribute.

653	Our observations warn of potential interpretation pitfalls in applying default attributes to					
654	real seismic data, especially when using geometrical attributes and others that are windows					
655	dependent. We offered a simplified workflow for geoscientists to understand and address these					
656	concerns depending on their available data.					
657	The importance of using synthetic data to reduce uncertainty is proved. This data allows					
658	essaying multiple scenarios to provide tools that serve geoscientists that face different kinds of					
659	datasets around the world and help reduce uncertainty by applying best practices in seismic					
660	interpretation, especially in channel deposit settings.					
661	Best practices in interpretation should be documented more often to better address					
662	uncertainty and optimize reservoir characterization.					
663	Data Availability					
664	Dataset presented at this article can be found at					
665	https://data.nzpam.govt.nz/GOLD/system/mainframe.asp, an open-source online data repository					
665 666	https://data.nzpam.govt.nz/GOLD/system/mainframe.asp, an open-source online data repository hosted at New Zealand and Petroleum Minerals. Synthetic data and model need to be requested					
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665 666 667 668 669	https://data.nzpam.govt.nz/GOLD/system/mainframe.asp, an open-source online data repository hosted at New Zealand and Petroleum Minerals. Synthetic data and model need to be requested to authors. Acknowledgments We thank SLB and AASPI (University of Oklahoma) for the software used for attribute					
665 666 667 668 669 670	https://data.nzpam.govt.nz/GOLD/system/mainframe.asp, an open-source online data repository hosted at New Zealand and Petroleum Minerals. Synthetic data and model need to be requested to authors. Acknowledgments We thank SLB and AASPI (University of Oklahoma) for the software used for attribute and noise computation and visualization; and the Chile Slope Systems Consortium (and					
665 666 667 668 669 670 671	https://data.nzpam.govt.nz/GOLD/system/mainframe.asp, an open-source online data repository hosted at New Zealand and Petroleum Minerals. Synthetic data and model need to be requested to authors. Acknowledgments We thank SLB and AASPI (University of Oklahoma) for the software used for attribute and noise computation and visualization; and the Chile Slope Systems Consortium (and sponsors) for providing the synthetic data and 3D model. Thanks to NZPM for providing the					
665 666 667 668 669 670 671 672	https://data.nzpam.govt.nz/GOLD/system/mainframe.asp, an open-source online data repository hosted at New Zealand and Petroleum Minerals. Synthetic data and model need to be requested to authors. Acknowledgments We thank SLB and AASPI (University of Oklahoma) for the software used for attribute and noise computation and visualization; and the Chile Slope Systems Consortium (and sponsors) for providing the synthetic data and 3D model. Thanks to NZPM for providing the Pipeline 3D dataset. Additionally, thanks to the AASPI group at OU and SEG organization for					
665 666 667 668 669 670 671 672 673	https://data.nzpam.govt.nz/GOLD/system/mainframe.asp, an open-source online data repository hosted at New Zealand and Petroleum Minerals. Synthetic data and model need to be requested to authors. Acknowledgments We thank SLB and AASPI (University of Oklahoma) for the software used for attribute and noise computation and visualization; and the Chile Slope Systems Consortium (and sponsors) for providing the synthetic data and 3D model. Thanks to NZPM for providing the Pipeline 3D dataset. Additionally, thanks to the AASPI group at OU and SEG organization for their financial support. Special appreciation to Thang Ha, David Lubo-Robles, Emily Jackson,					

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Figure 1: Stratigraphic hierarchy: from a single element to complex sets. A vertical slice within 812 the Pipeline 3D dataset offshore New Zealand shows the seismic appearance of different 813 architectural hierarchies in deepwater channels, with corresponding cartoons below. The smallest 814 architecture (4th to 5th order) is the channel element (box 1 in green). The second hierarchy (6th 815 order) occurs when the channel elements stack together, forming a channel complex (box 2 in 816 blue). The higher-order hierarchy (7<sup>th</sup> order and higher) occurs with the amalgamation of channel 817 complexes, developing a channel complex set (box 3 in magenta). The color legend indicates the 818 819 distinct facies that commonly occur within each element of each architecture. Measuring the sizes 820 of each architectural element indicated on the right as well as their hierarchy, provides key insight 821 into the underlying depositional processes as well as a prediction of the more common lithologies. Hierarchies mentioned follow Pickering and Cantalejo (2015) classification. 822

823 Figure 2. A visual guide showing the steps to convert outcrop measurements to a synthetic model. 824 (A) Location and exposure of the outcropping deepwater channels at Laguna Figueroa. Paleoflow is from North to South (obliquely and to the right into the outcrop at this location) (B) Conceptual 825 diagram of the Upper and Lower Figueroa outcrops showing channel elements, complexes and 826 complex sets. The red line indicates the outcrop profile. Left of the line is into the outcrop face, 827 and the right has been eroded away. (C) geocellular model using the constraints from (B) 828 augmented by facies and corresponding rock properties, including acoustic impedances from 829 830 Shallow Offshore West African modeled rock properties (Stright et al., 2014). (D) The Ormsby wavelet and a representative vertical slice through the 3D synthetic seismic data volume generated 831 from the model shown in (C). Courtesy of Teresa Langenkamp and Lisa Stright. 832

Figure 3. Explanation of stratigraphic event mixing or vertical smearing of stratigraphic features. 833 (A) Vertical resolution of channel elements related to different peak frequencies (modified after 834 Nielson, 2011). (B) Vertical slice through the 3D Pipeline 3D offshore New Zealand seismic 835 survey showing a channel complex and a Horizon used for interpretation at different stratigraphic 836 levels indicated by yellow arrows. The dominant frequency at this level is 40 Hz giving a dominant 837 period of 25 ms. (C) Stratal slices at approximately 25 ms intervals through the coherence volume 838 computed using a  $\pm 20$  ms analysis window show "stratigraphic" mixing by the seismic wavelet. 839 Note that the relatively straight channel form at 2050 ms can be seen at 2080 ms and other 840

stratigraphic levels (1,2,3 from deeper to shallower) where other channel forms (green arrow) appear causing interference. The cause of this mixing could be 1) mixing of reflectivity by the 25ms dominant period seismic wavelet, 2) mixing of discontinuities through the 40 ms coherence computation, 3) shifting of the basal channel element thalweg due to compensational style as you move up or 4) differential compaction over deeper discontinuities between the floodplain and the channel element fill.

**Figure 4.** Workflow of the study. Four different parameters were evaluated in the sensitivity analysis: 1) the effect of frequency content, 2) the impact on the choice of the seismic attribute, 3) the analysis window effect, and 4) the sensitivity to band-limited random noise addition. The best cases were selected to be shown in each case, and analysis was performed by comparing them with the original 3D geological model derived from the Laguna Figueroa Deepwater outcrop. This resulted in 4 cases for analysis, ultimately leading to a workflow and documentation of best practices in channel architecture interpretation.

Figure 5. A representative analysis window used in attribute calculations. Signals are sampled at discrete points, not continuous recordings. Therefore, each seismic trace will contain as many samples as the sample rate allows. For the example shown, if we consider the dominant period (distance between two peaks), our analysis window will contain 20 samples, equivalent to  $\sim 20$ msec. Other examples of analysis windows are depicted, including the smallest possible equivalent to the sample increment, in this case, 1 msec.

Figure 6. A representative vertical slice (224) and time slices at t=-191ms) through the (A) 3D model that shows (B) 15 Hz data, (C) 60 Hz data, (D) 90 Hz data, and (E) 180 Hz data. Notice the improvement in the channel architecture's detail with the increase in the dominant frequency and corresponding spectral bandwidth, being able to interpret complexes from figure B on.

**Figure 7**. Results on the attribute sensitivity analysis to thickness and extent of architectural elements using a default analysis window. We present results on the 30 Hz dataset, analogous to vintage seismic data (to the left), and the 60Hz dataset, representing modern seismic data (to the right). Each attribute is presented for both frequencies in a representative inline (224) and time slices (-280 ms and -191ms) as follows: (A) RMS (Root Mean Square) amplitude for 30 Hz (B) coherence (Sobel filter) for 30Hz, (C) Instantaneous frequency for 30 Hz, (D) RMS amplitude for 60 Hz (E) coherence (Sobel filter) for 60Hz, (F) Instantaneous frequency for 60 Hz. Figure 8. RMS amplitude sensitivity analysis to evaluate the effect of the window size W on the vertical smearing of the different channel elements in the model. (A) 15 Hz with W=2ms (B)50 Hz with W=50ms (C) 30Hz with W=2ms (D) 50 Hz with W=50ms (E) 180 Hz with W=2ms (F) 180 Hz with W=50 ms. High-frequency data with a small analysis window provides the most suitable representation of the true model. In RMS amplitude, the most accurate facies depiction is given by high frequencies and small window combination, imaging channel element base, and inner and outer levee facies.

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Figure 9. Sobel filter coherence attribute sensitivity analysis to evaluate the effect of the window size W on the vertical smearing and displacement of the different architectural elements in the model. (A) 15 Hz with W=2ms (B)30 Hz with W=2ms (C) 60Hz with W=2ms (D) 60 Hz with W=50ms. Notice how at a higher dominant frequency (60Hz) and smaller window of analysis, there is better detection of the channel element edges and channel element fill. In contrast, at lower frequencies and or larger windows of analysis, there is interference from deeper channel elements, as in the case presented on Figure 3.

886 Figure 10. Sensitivity analysis of the effect of random noise in channel architecture interpretation. Results shown refer to the 60 Hz dominant frequency seismic volume (which would be analog to 887 888 real datasets). (A) 5% band-limited random noise added evaluated in RMS, Coherence, and Instantaneous frequency attributes. Interpretation of prominent features like channel complexes is 889 890 only possible. (B) 5% band-limited random noise added applied to RMS, coherence, and instantaneous frequency attributes. Notice how the increase in noise is detrimental in the channel 891 architecture interpretation, especially when using coherence. This may be due to the sensitivity 892 that small windows have on high noise content. 893

Figure 11. Generalized workflow for a geoscientist to avoid pitfalls in interpretation by gettingoptimized results according to their dataset.

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#### LIST OF TABLES

Table 1. Tuning thicknesses for shallow and deep elements in each synthetic seismic volume.
Modified from Langenkamp (2021).

**Table 2**. Selected seismic attributes to perform the sensitivity analysis showing a representative

900 image, feature measurement, and use in channel element interpretation.

901 **Table 3**. Common pitfalls in seismic interpretation and how to avoid them.

# APPENDIX

Dominant			Low-cut	Low-Pass	High-Pass	High-cut
Frequency	Length	Sample	Frequency	Frequency	Frequency	Frequency
(Hz)	(ms)	rate (ms)	(Hz)	(Hz)	(Hz)	(Hz)
15	200	1	1	3	23	35
30	200	1	2	6	45	70
60	100	1	4	12	90	140
90	100	1	6	18	135	210
180	26	1	10	30	225	350

A- Specs for each synthetic volume

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